Far-infrared properties of phyllosilicates & carbonates

Lorentz Center Leiden, 3.3.2011 Workshop "Herschel and the Characteristics of Dust in Galaxies"

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What I would wish to show you:

Interesting FIR lab spectra... ... compared to nice new spectra...



& cool crystals...

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But this cannot be done yet!

... since there are still problems with Herschel-PACS spectroscopy:

- Removal of molecular lines
- Reliable determination of background slopes of spectra
- Especially important / problematic for broad bands



Herschel-PACS spectroscopy vs. imaging

- Herschel-PACS imaging has already revealed a lot about dusty stellar environments
- <u>But</u>: PACS' spectroscopy of dust still causes some headache.







Kerschbaum et al. 2010

Fig. 1. *Herschel*-PACS scan maps (*left to right*: at 70 μ m, 160 μ m, two colour composite) of AQ And, U Ant, and TT Cyg (*top to bottom*).

What <u>can</u> be done:

Discussion of FIR properties of selected dust species as measured in the Jena lab



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Dust features in the far-infrared

- Most (diagnostically relevant) dust features are in the MIR: 8-40 μm
- Mineral spectroscopy also focuses on the MIR
- But there do exist FIR dust features
- Among the carriers: heavy ions / larger parts of lattices

$$\lambda_{vib} \propto \sqrt{\frac{\mu}{k}}$$

- E.g., ion groups or lattice sheets in vibration
- To actually <u>see</u> astronomical FIR, need additionally: cool dust

Carbonates, phyllosilicates and the "landscape" of dust



I. Phyllosilicates





Up to 10% of phyllosilicates In HD 142527? (Bowey&Adamson 2002)

Why to study phyllosilicates

- Have pronounced bands in the FIR
- Form by aqueous alteration of pyroxene and olivine minerals (e.g. enstatite → talc; forsterite → talc)
- Important constituents of primitive meteorites (CM, CI carbonaceous chondrites)
- Occur in IDPs (Dorschner et al. 1978; Knacke et al. 1980; Sandford et al. 1985)
- Important for astrophysics as a potential water reservoir in space
- Existence in protoplanetary disks was proposed (Malfait et al. 1999 based on Koike et al. 1982) (see below: broad shoulder from ~85-125µm)





Possibly, water on <u>Earth</u> actually came from hydrous phyllosilicates contained in <u>meteorites</u> in the early solar system

The presence of phyllosilicates in protoplanetary disks would be important e.g. for "water transport"

Why to do low temperature spectroscopy

- Absorption/emission behavior of crystalline matter depends on temperature (anharmonicity of the potential of binding of atoms in the crystal structure)
- Decreasing temperature \rightarrow increasing energy hv of photons for activation of next vibrational state \rightarrow bands shift to shorter λ 's
- In some cases: FIR bands = ",thermometers" which is interesting





Different Kinds of Silicates

- Basic units of all silicates: [SiO₄]⁴⁻ -tetrahedra
- One oxygen atom can be a part of another tetrahedron at the same time
- Many different structures can be built: Isolated tetraedra (olivines), chains (pyroxenes), sheets ($\varphi U \lambda \lambda O V = leaf$) (phyllosilicates)



isolated



layers

<u>Phyllosilicates:</u> 3 oxygen atoms of one tetrahedron are a part of other tetrahedra in one layer – built up layerwise, between the layer: metal cations (K⁺, Na⁺, Ca²⁺, Al³⁺), OH⁻ and H₂O (\rightarrow hence: "hydrous silicates")

Experimental Setup (Jena lab)



Experimental Details (I)

- Measurements taken at 300K, 200K, 100K and 10K
- Sample preparation: grinding in a mortar and/or ball mill and sedimentation in acetone to grain sizes ${<}2\mu m$

Phyllosilicate	Formula	group
Chamosite	$Fe_{3,55}Al_{1,88}[(A1,Si_3)O_{10}(OH)_8]$	Chlorite
Talc	Mg _{3,33} Fe _{0,1} [Si ₄ O ₁₀ (OH) ₂]	Talc-Pyrophyllite
Montmorillonite	$\begin{array}{l} Al_{1,5}Mg_{0,25}Fe_{0,17} \texttt{[Si_4O_{10}(OH)_2](Na,K) \cdot I,2} \\ H_2O \end{array}$	Clay
Picrolite	Mg _{5,84} Fe _{0,17} [Si ₄ O ₁₀ (OH) ₈]	Serpentine
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Experimental Details (II)

All samples were pressed to polyethylene pellets (m=200mg, d=13mm)

Embedding ratios (sample mass : PE mass)

	17-67µm	67-200µm	>200µm
Phyllosilicates	1:100	1:100/1:10	1:10



Measured transmittance spectra



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From transmittance to MAC spectra

Deriving κ from the transmittance:

$$\kappa = -\frac{A}{m_{emb}} \cdot \ln(T)$$

Fringe correction by averaging adjacent data points



Fig. 6. Mass absorption coefficient spectra of the chamosite sample in the FIR range. The spectra are a combinations of measurements at two different column densities (see text). The fringes in the original data have been removed by a sliding average method used up to a wavelength of about 200 μ m and by Lorentzian fitting for the 277 μ m band. The structure at about 130 μ m is due to imperfect compensation of a PE absorption band.

Further findings at low temperatures

- Weak bands become stronger and sharper
- ➢ "Hidden" bands can be discovered
- > For montmorillonite, the FIR gradient changes from ~ I/λ to I/λ^2 upon cooling



Comparison with ISO-LWS data



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Summary on Phyllosilicates

- Montmorillonite does not only show a broad, 85-125µm FIR band!
- •We discovered a previously unknown band of **talc** at 98.5µm (300K value) and several "hidden" bands of **montmorillonite**
- Most bands shift to smaller λ 's at decreasing temperatures \rightarrow "thermometer" effect
- We can't confirm the existence of phyllosilicates outside our solar system yet: Spectra of protoplanetary dust disks are too inconclusive
- Search for talc / montmorillonite bands: must focs on **relatively sharp rather than broad bands**

II. Carbonates



- Cold dust in Planetary Nebulae
- Environments of protostars
- Mars' atmosphere: CaCO₃, FeCO₃ aerosols
- IDPs (e.g. Sandford 1986): contain solid solutions of $MgCO_3$ and $FeCO_3$
- Comets, e.g. Tempel 1

<u>Atmosphere</u> of Mars: may contain carbonate dust (ISO-SWS absorption spectra)



Examples of potential carbonate bands



ISO-SWS&LWS

Planetary Nebulae NGC 6302 & 6537: FIR bands at ~65µm and ~90µm ascribed to CaMg[CO₃]₂ and CaCO₃ dust by Kemper et al. (2002), Nature 415, 295



Comet 9P/Tempel 1: NIR bands around 6.8µm, detected by Spitzer in "Deep Impact" ejecta, ascribed to MgCO₃ and FeCO₃ by Lisse et al. (2006), Science 313, 635

Small particle spectra (I)

- NGC 6302 (cont. subtracted)
- --- calcite, CDE, 30K (abs.eff. x Planck function)
- calcite, CDE, 60K (abs.eff. x Planck function)
- dolomite, CDE, 30K (abs.eff. x Planck function)
- ----- dolomite, CDE, 60K (abs.eff. x Planck function)



Posch, Mutschke et al., ApJ 668, 993

- Key results in comparison to stardust spectra:
 - Dolomite has strong ~65µm emission band.
 Profile strongly dependent on shape and temperature.
 In no case, ~65µm band in NGC 6302
 explainable by dolomite dust

Calcite has strong ~90µm emission band.
 For T<45K, all other calcite bands
 suppressed & observed band profile compatible
 with irregularly shaped CaCO₃ grains



Small particle spectra (II)

Powder spectra were also measured (at room temperature)

Results of measurements agree well with CDE calculations



NB: Presence of large grains in unsedimented powder New measurements lead to different band positions and relative strengths Compatible with ISO observations of NGC 6302, but setting more narrow constraints on possible temperatures (below 50K)

Carbonate data: available online

- Optical constants for calcite and dolomite available at: <u>http://www.astro.uni-jena.de/Laboratory/OCDB</u>
- Including plots of $n(\lambda)$, $k(\lambda)$ and $Q_{abs}(\lambda)$







Places of work & Acknowledgements



Jena

Space (Herschel)



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Thanks to:

- Gabriele Born
- Thomas Henning
- Walter Teuschel

Vienna